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CORA – a dedicated device for carbon dioxide monitoring in cave environments

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Abstract:

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High resolution time-series of cave CO₂ fluctuations are increasingly demanded to quantify calcite precipitation processes. CORA, an energy-efficient NDIR-device, has been specifically developed for the long-term monitoring of carbon dioxide in remote cave environments. To allow comparison between different cave sites, changes in air pressure and temperature are compensated for using dedicated probes. Laboratory experiments demonstrate that CORA's precision is adapted to the analysis of spatially and temporally variable CO₂ regimes and therefore suitable for a large number of applications. Data obtained with 12 independently calibrated instruments are reproducible within 3% (1 σ). The two-point calibration function is validated using certified reference gases for an overall accuracy determined at ± 50 ppmv, resp. $\leq \pm 3\%$. Although no significant drift was observed during the test phase a periodic recalibration is recommended. Initial results from field studies show that continuous monitoring over a full annual cycle is possible at 4 hour intervals without specific maintenance.

Keywords: cave; carbon dioxide (CO₂); monitoring; climate; speleothem

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INTRODUCTION

Carbon dioxide is a key component for karstification, namely the dissolution and precipitation of CaCO₃ which affects most carbonate rocks. The initiation and development of karst conduits results in complex drainage patterns controlling the subsurface hydrology and meteorology (Ford & Williams, 2007). Because of the transient transport processes and spatial heterogeneity inherent to these karst systems, the subsurface PCO₂ is typically increased as compared to the external atmosphere and may vary both in space and time (e.g., Bourges et al., 2001; Fernandez-Cortes et al., 2006, 2009; Baldini et al., 2006).

Previous studies have identified three potential sources for carbon dioxide in caves (James, 1977): 1) soil respiration; 2) oxidation of organics transported into the cave and, 3) a deep-seated, hydrothermal origin. Depending on production rates and ventilation regimes cave air CO₂ concentrations may vary by several orders of magnitude. In vadose cave environments, CO₂ concentrations typically range between values close to the external atmosphere and about 4% (e.g. Troester & White, 1984; Ek &

Gewelt, 1985; Bourges et al., 2006; Batiot-Guilhe et al., 2007; Liñán et al., 2008) but extreme values of up to 18% have occasionally been reported (Halliday, 1984). Studying cave PCO₂ fluctuations is therefore fundamental to quantify dissolution and precipitation rates of calcium carbonate and assess the contribution of karst processes onto the global carbon cycle. Moreover, elevated cave air CO₂ concentrations may be an indicator of pollution in a cave or its catchment area (e.g. Deboz et al., 2002), which may eventually affect groundwater resources.

The interest in the spatial distribution and the temporal evolution of cave PCO₂ has significantly risen with the successful use of speleothems as terrestrial paleoclimate archives (e.g. review by Fairchild et al., 2006). Banner et al. (2007) demonstrated that changes in the cave PCO₂ modify stalagmite growth rates and thus impinge on paleoclimate proxies, including stable isotopes (Spötl et al., 2005; Matthey et al., 2008, Frisia et al., 2011) and trace elements (Treble, 2005; Johnson, 2006; Wong et al., 2011). So far, cave CO₂ data were however, often limited to discontinuous measurements and only during the recent years, with the onset of non-dispersive infrared (NDIR) measurement techniques, continuous monitoring of cave CO₂ has become more frequent (e.g. Baldini et al., 2008; Kowalczyk & Froelich, 2009; Sherwin & Baldini, 2011).

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Unfortunately, commercially available instruments rarely suit the specific requirements for cave monitoring and many systems face severe limitations due to power supply and/or electronics failure due to elevated condensation rates. In absence of coupled probes, correction for pressure and temperature changes is often hindered, making a comparison between monitoring sites difficult. Here we present a new energy-efficient NDIR-device for the long-term monitoring of CO₂ concentrations in cave environments. This hand-held instrument was specifically developed for exigent monitoring conditions (e.g. no grid supply, difficult access, low temperature, high humidity, muddy conditions) and successfully tested for several months in an alpine cave. Versions of this instrument are currently operating at cave sites across Europe (Bärenhöhle, Katerloch, Obir, and Spannagelhöhle, Austria; Milchbachsystem and Milandre cave, Switzerland, and Conturines cave, Italy) as well as in Australia (Wombeyan cavern).

NDIR GAS ANALYSIS – METHODOLOGICAL BACKGROUND

It has been empirically demonstrated that various gases absorb radiation in the infrared portion of the electromagnetic spectrum due to intramolecular vibrations. According to the wavelength of the incident radiation (typically between 2 and 20 μm) and depending on the chemical bond structure of the target molecules, a series of absorption peaks will be observed. The non-dispersive measurement technique uses fixed band-pass filters to identify characteristic absorption lines over a selected wavelength range (e.g. Stuart, 1996). The intensity of IR light reaching the detector is inversely related to the concentration of the target gas in the sample chamber. Following the Beer-Lambert law (Eq. 1), the concentration of the target gas is given by:

$$I = I_0 e^{-\alpha LC} \quad (1)$$

where I is the intensity measured in the target gas [V]; I_0 is the intensity measured in absence of target gas (zero gas) [V]; α is a gas-specific absorption coefficient [$\text{m}^3 \cdot \text{mol}^{-1} \cdot \text{cm}^{-1}$]; L is the length of the optical path [cm] and C is the concentration of the target gas.

This idealised expression faces issues related to the mechanical and optical alignment of the sensor components, lamp intensities and spectral output. To compensate for non-linearity effects Equation (1) is modified by introducing two empirical parameters (e2v, 2007):

$$I = I_0 \exp(-aC^n) \quad (2)$$

where a and n are specific to a particular sensor design and determined for individual concentration ranges of a given target gas.

The linearization coefficients a and n , which can be determined empirically from measurements in reference gases at different concentrations (e2v, 2007) are typically provided by the manufacturer. Using a dual channel NDIR sensor allows comparing

the detected signal with a reference signal filtered at a band width displaying no absorption lines for the specified gas. The change in output voltage measured in the presence of the target gas provides the normalized absorbance (Abs):

$$\text{Abs} = 1 - \text{Act} / (\text{Zero} * \text{Ref}) \quad (3)$$

where Act = active signal; Ref = reference signal; Zero = Act / Ref in zero gas.

Combining Equation (2) and (3) and introducing a gas span accounting for the proportion of absorbing radiation gives:

$$1 - \text{Act} / (\text{Zero} * \text{Ref}) = \text{Span} (1 - \exp(-aC^n)) \quad (4)$$

From equation (4), we derive the target gas concentration (C):

$$C = \left[\frac{\ln\left(1 - \frac{\text{Abs}}{\text{Span}}\right)}{-a} \right]^{\frac{1}{n}} \quad (5)$$

Changes in the ambient temperature and pressure can be compensated for by using the ideal gas law, providing the instrument is equipped with independent temperature and pressure probes. Self-heating effects due to the light source must be compensated by introducing a linear correction for Abs and Span using a separate temperature sensor mounted close to the gas sensor. For further details the reader should consult the manufacturers' application notes (e.g. e2v, 2007; Alphasense, 2009).

CORA – A CO₂ RECORDING APPARATUS

The CORA^{1*} monitoring device comprises a CO₂, a pressure, and a temperature probe (Table 1). Recorded data are stored on a 32 Mbit internal memory and can be transferred to an external SD-card peripheral or downloaded onto a notebook via a RS485 serial port. Thanks to its low-power consumption the device allows monitoring at 4 hour intervals for at least one full year without maintenance. Data exchange with the instrument software is achieved through a Membrane Keypad allowing programming the measurement time interval (5 min to 48h), to start monitoring sequences and to download data.

CORA was specifically developed to resist dirt and water splashes (IP52) typical of cave environments. The device consists of a 15 x 9 x 5 cm handheld apparatus with a digital display. Six internal alkaline AA cells (9V; 3.4 Ah) supply sufficient energy for one month autonomy. The instrument weighs 590 g and is packed in a waterproofed Otter Case for transport (Fig. 1). For monitoring periods exceeding one month, the instrument can be coupled to an external battery pack (12V; 22 Ah; Fig. 2). The total weight of CORA including external batteries is slightly less than 3 kg.

¹Cora, also known as Persephone, reigned the underworld together with Hades in the Greek mythology.



Fig. 1. The CORA device used for monitoring CO₂ concentrations in cave air. The yellow box underneath the device holds eight 1.5 V batteries (D cells).

User interface

CORA’s display provides information about the instrument status and the last recorded measurement. Navigation is achieved via the membrane keypad using up and down arrows combined with “OK” and “Cancel” touches. The software menu is structured into six topics each comprising a list of relevant options: 1) measurement, 2) time interval, 3) export data, 4) settings, 5) switch off, and 6) info.

The user can choose between performing a single measurement and starting a monitoring period. The measurement time interval can be defined separately for the carbon dioxide and temperature/pressure. This option allows differentiated monitoring strategies taking into account the higher power consumption of the CO₂ sensor, eg. 4 h for the carbon dioxide but 10 min for the other two parameters.

The settings allow selecting different display characteristics, including the contrast and backlight intensity. The instrument’s date is defined manually to account for monitoring projects in different time zones. The user has also the opportunity to choose between three different languages (English, German and French); further languages will be added in the future.

The calibration menu is password-protected and an administrator can change the parameters

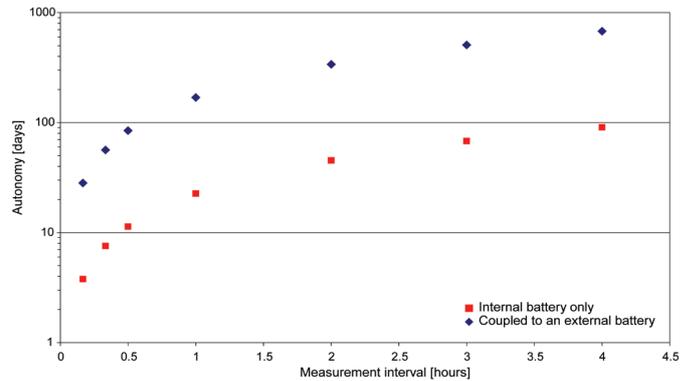


Fig. 2. Expected power autonomy of CORA according to the measured time interval. The values represent a conservative estimate valid for cold alpine cave environments.

defining the CO₂ calibration curve including empirical determinations of the linearization coefficients and calibration parameters for the temperature and pressure probes.

Technical - Sensor(s) – Electronics

Changes in the atmospheric PCO₂ are measured with an e2v dual channel NDIR sensor by comparing the output signal of the detector filtered to a CO₂ absorption band centred at 4.2 μm with a reference signal (λ= ~4.0μm). The IR source is provided by a tungsten filament hosted in a glass envelope while the air sample to be analysed enters the optical cell through a particulate filter by molecular diffusion (e2v, 2008). The measured signal is proportional to the molecular density of CO₂ present in the ambient air. Compensation of temperature and light source drifts is achieved through a reference channel warranting the long-term stability of the output signal (e2v, 2008).

CORA relies on a 6V DC power supply which is switched on at regular time intervals for single measurements. To offset thermal background signals due to changes in light levels the IR source is modulated at 2Hz. The resulting signal measured at the pyroelectric detector is sinusoidal. The peak-to-peak AC voltage is proportional to the incident IR allowing for synchronous detection techniques. Every CO₂ measurement requires typically 50 mA over an

Table 1. CORA sensor specifications

Sensor	NDIR-CO ₂	Temperature	Air Pressure
Manufacture	e2v	IST AG	Freescale Semiconductor
model	IR11GJ	Pt 1000 Class A	MPXH6115A6U
Power requirement	9-12 V, 80 mA	0.3 - 2 mA	3V, 4 mA
Measurement range	0-2% vol. CO ₂	-50 - +150°C	15 - 115 kPa
Response time (signal averaging)	180 sec	ca. 20 sec	1 ms
Warm-up time	90 sec	-	20 ms
Accuracy	± 3% of reading	± 0.2°C	± 1.5% of range
Long-term drift *	0.05%vol. /month	0.04 % / 1000 h	± 0.25 % / 1000 cycles

* according to manufactures notes

Table 2. Reproducibility of CO₂ measurements using 12 different CORA devices and a certified reference gas (7384±147 ppmv_{satp}); n.d.: not determined.

CORA #	Normalised absorbance	Calibrated [CO ₂]	Number of meas.
2	0.230060 ± 0.000839	7267 ± 49	3
3	0.239571 ± 0.001962	7494 ± 174	8
4	0.219034 ± 0.003982	7248 ± 78	2
10	0.125561 ± 0.001061	7442 ± 122	9
11	0.097857 ± 0.000078	7644 ± 35	3
12	n.d. n.d.	n.d. n.d.	
13	0.104523 ± 0.000437	7766 ± 105	6
14	n.d. n.d.	n.d. n.d.	
21	0.301808 ± 0.002277	7481 ± 16	3
22	0.298984 ± 0.002963	7549 ± 113	3
23	0.305982 ± 0.002736	7488 ± 77	2
24	0.302495 ± 0.000931	7356 ± 15	3
25	0.309169 ± 0.002164	7515 ± 101	3
26	0.324852 ± 0.001300	7842 ± 42	3
		7508 ± 179 ppmv satp	

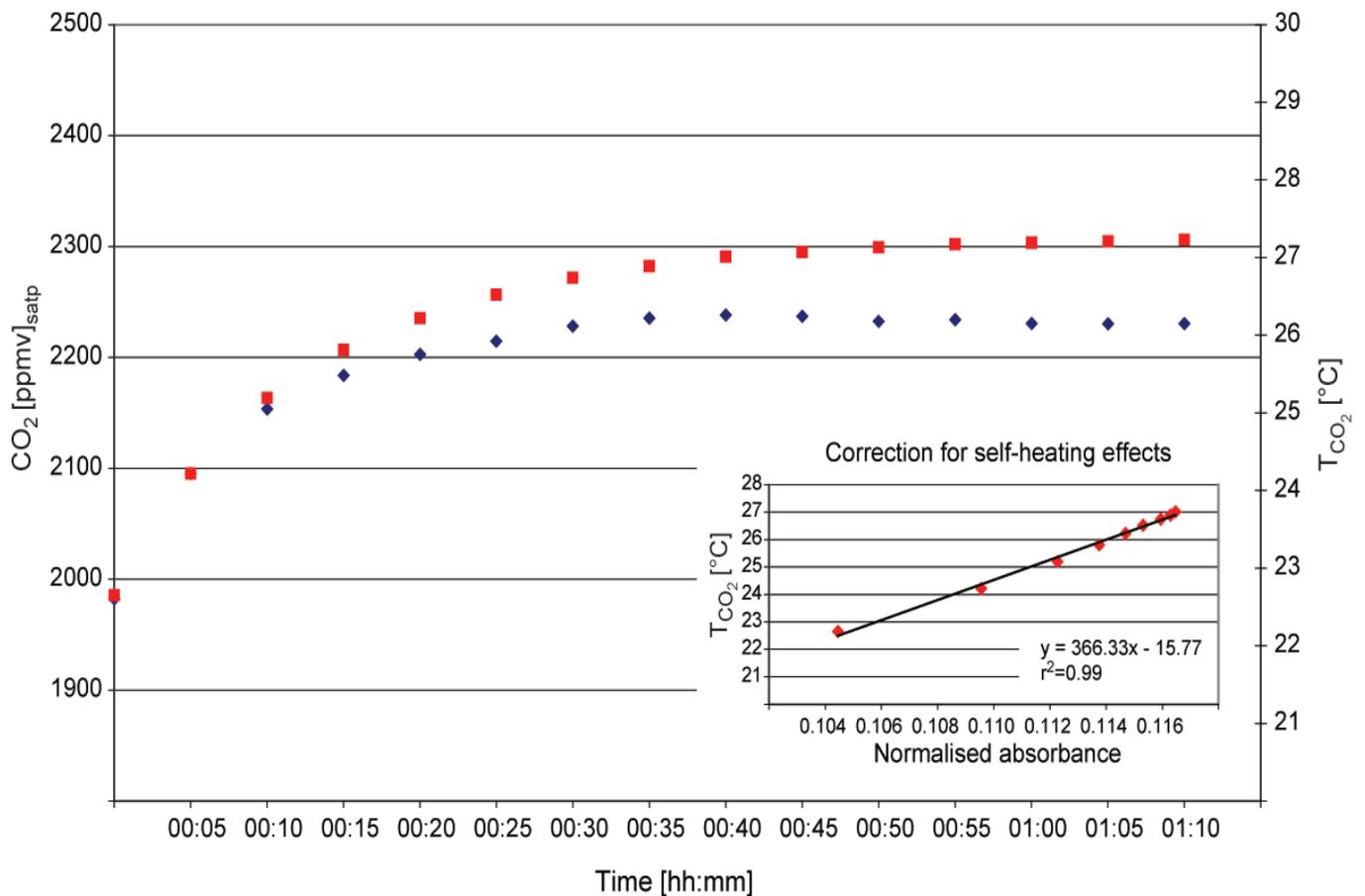


Fig. 3. Temperature dependency of CO₂ readings at 5 min intervals and constant PCO₂. Blue diamonds: measured [CO₂] values; red squares: temperature of the NDIR sensor (TCO₂). The measurement at t₀ corresponds to the actual calibrated measurement. Self-heating associated to short measurement intervals can be corrected for (inset).

acquisition window lasting 270 s. During the first 90 s the sensor is preheated and the data are disregarded. Measurements performed during the following 180 s are averaged to the final result using a low-pass filter. Digital signal processing provides the temperature- and pressure-corrected CO₂ mole fraction in ambient air [μmole/mole].

AIR TEMPERATURE AND PRESSURE COMPENSATION

To allow a comparison of carbon dioxide measurements between different cave sites as well as in large cave systems showing a range in air pressure and temperature, the measured concentration of CO₂ is compensated for atmospheric changes:

$$[CO_2]_{satp} = [CO_2]_{meas.} \frac{1013.25 * T_{meas.}}{288.15 * P_{meas.}} \quad (6)$$

where [CO₂]_{satp} is the CO₂ molar concentration [ppm] in the standard atmosphere at constant temperature and pressure (satp; 25°C, 1013.25 hPa); [CO₂]_{meas.} is the measured CO₂ molar concentration; T_{meas.} is the cave air temperature [K] and P_{meas.} is the barometric pressure [hPa].

TEMPERATURE DEPENDENCY – SELF-HEATING

Temperature effects due to self-heating of the NDIR sensor are compensated for at each acquisition window using an internal thermistor. As full thermal equilibrium of the sensor is only reached after ca.

30 min (Fig. 3) analysing transient thermal regimes may require an additional correction to ensure a high accuracy. This may be particularly relevant in the entrance zone of warm cave systems where the heat exchange between the sensor and the subsurface atmosphere is slower than in colder cave environments. Overall, it is recommended to keep the measurement interval at minimum 15 minutes.

CORA PERFORMANCES

Precision and long-term drift

The short-term precision of CORA measurements was tested at 2 s intervals over 900 s. After correcting for temperature effects due to sensor heating, the relative standard deviation of the normalised absorbance measured in air at constant PCO₂ (i.e. 100% N₂) reaches ±0.02%. Similar results (i.e. ±0.03%) were obtained from 153 individual measurements acquired over 77 h at 20 min intervals (CORA #13).

The long-term precision of the instrument was determined at daily intervals over a period of 8 months with a ~2000 ppmv CO₂ reference gas (1832±40 ppm satp) involving a total of 138 individual measurements (CORA #2). Results shown in Fig. 4 suggest no significant drift during the observation period at an overall precision of the normalised absorbance of ±0.1% (1 sigma). In terms of the CO₂ concentration the latter corresponds to an absolute precision of 1829±29 ppmv, i.e. 1.6% of the measured value. This five times larger variance as compared to the short-term precision is primarily attributed to secondary effects associated with fluctuations of the battery voltage.

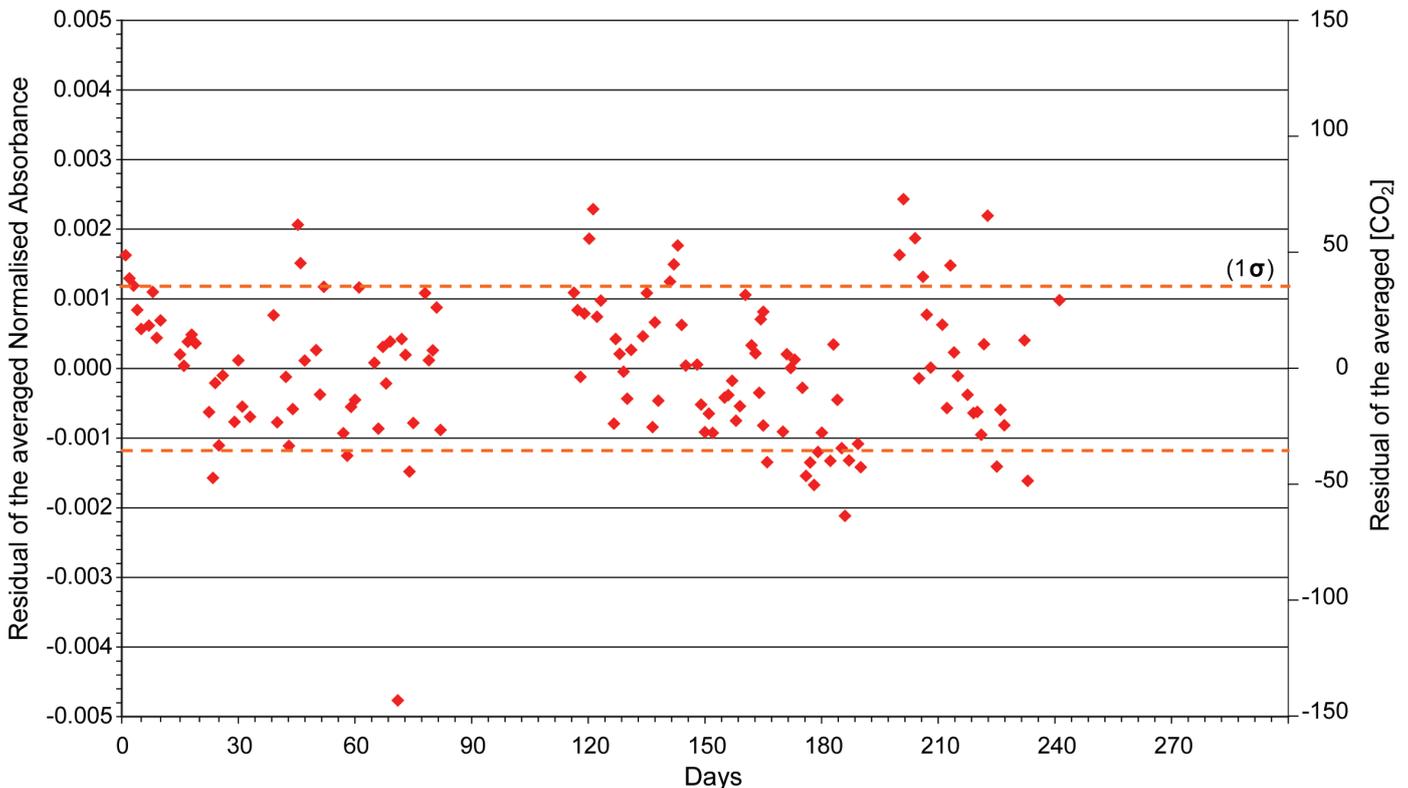


Fig. 4. Deviation from the mean normalised absorbance after correcting for temperature and pressure. The data was measured at roughly daily intervals using a 1832±40 ppmv_{satp} reference gas over a period of eight months (CORA-#3).

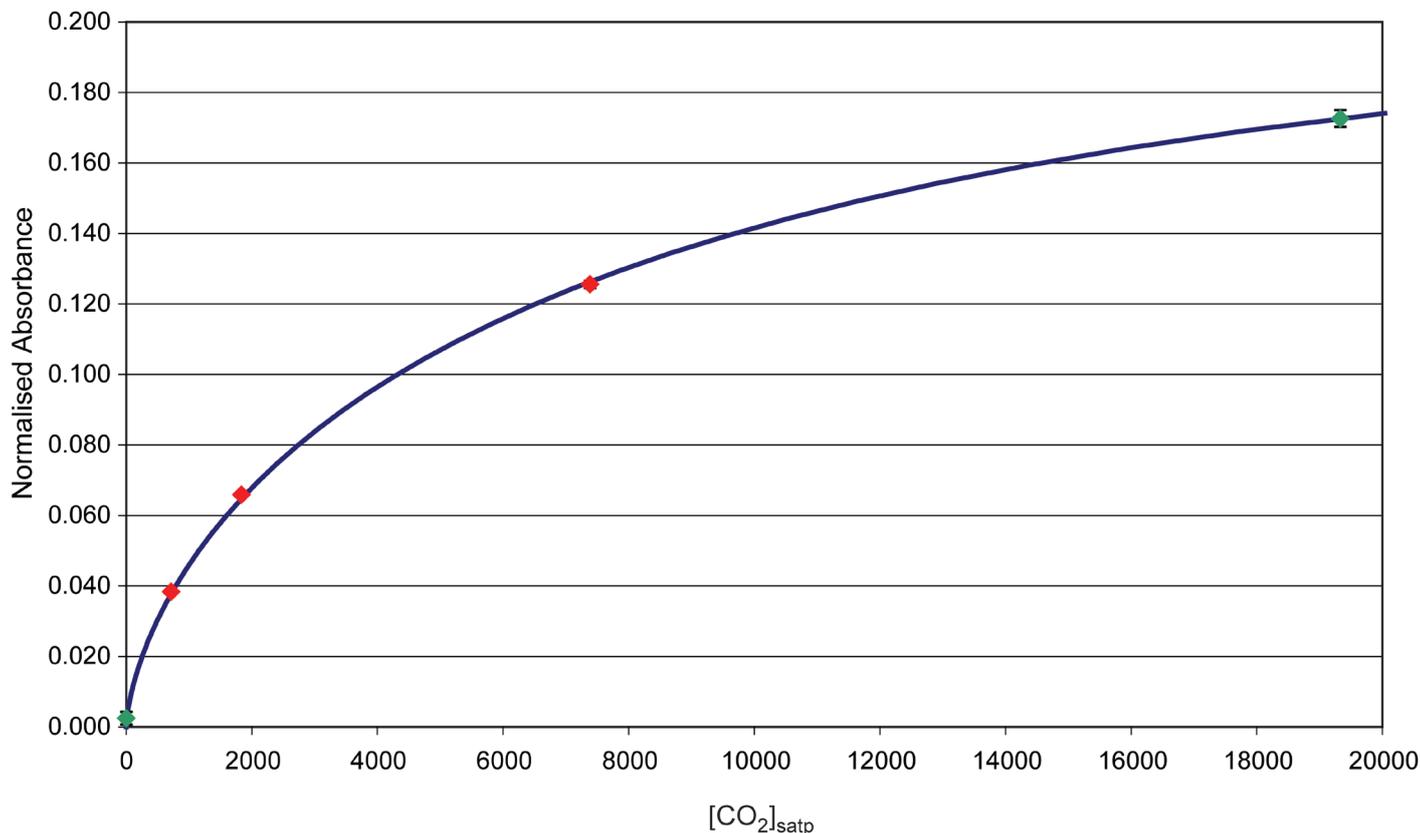


Fig. 5. [CO₂] values obtained using a 2-point calibration curve at 0 and 2% (green dots). Validation of CORA's accuracy is given by the measurement of three certified reference gases (red dots). Error bars are smaller than the dot size.

Accuracy

A 2-point calibration at 0 and 2% CO₂ allows converting the measured output signal into concentrations according to Equation (5). Fig. 5 shows an independent validation of this calibration curve by measuring the absorbance of three reference gases at ca. 800, 2000 and 8000 ppmv. All measurements were performed in lab conditions (i.e. dry gas samples) and compensated for temperature and pressure changes using the ideal gas law. The results point to a high accuracy of the instrumental output with actual values within 3% of the predicted concentration.

Reproducibility

Table 2 shows CO₂ concentrations measured using a certified reference gas (i.e. 7384±147 ppmv_{satp}) with 12 independently calibrated CORA devices. Owing to three different sensor generations, a difference is noticed in the normalised absorbance between CORA 2 to 4, 10 to 14 and 21 to 26, respectively. The calculated concentrations, however, are consistent with values varying between 7248 and 7842 ppmv and a mean of 7508 ± 179 ppmv. Taking into account the error margins all instruments reproduce within 3% (1s confidence interval) of the actual CO₂ concentration. The variable instrumental precision observed during this reproducibility test is most likely related to the small number of data available for some of the devices and is statistically not significant. Overall, the accuracy of the determined [CO₂] data is consistent with previous results.

DISCUSSION AND APPLICATION

The exponential nature of the Beer-Lambert law results in more precise measurements at low CO₂ concentrations. Still, empirical adjustments of linearization coefficients are necessary to achieve adequate accuracy. This requires calibration measurements to be carried out using certified reference gases under controlled atmospheric conditions. The validation of the resulting calibration curve is provided by independent measurements at known CO₂ concentrations.

Laboratory experiments have shown that the precision (±0.1% of the normalised absorbance, i.e. 1829±29 ppmv, Fig. 4) and the accuracy (±3%, Fig. 5) of CORA measurements are within the manufacturer's specifications for the NDIR sensor in use (e2v, 2008). Nonetheless, Fig. 4 underlines a discrepancy between the short and long-term precision which is attributed to fluctuations in the battery voltage. The latter is assumed to impinge on the rate of instrumental self-heating which is corrected for using an independent temperature probe mounted close to the CO₂ sensor. For long-term monitoring periods during which batteries are exchanged at regular intervals, the loss of precision has only a small effect. Over a single battery cycle, this may, however, result in a linear trend towards lower [CO₂]. Hardly noticeable at elevated concentrations, the drift becomes significant at CO₂ concentrations close to that of the external atmosphere and needs to be compensated for using a linear detrending function. Stabilisation of the

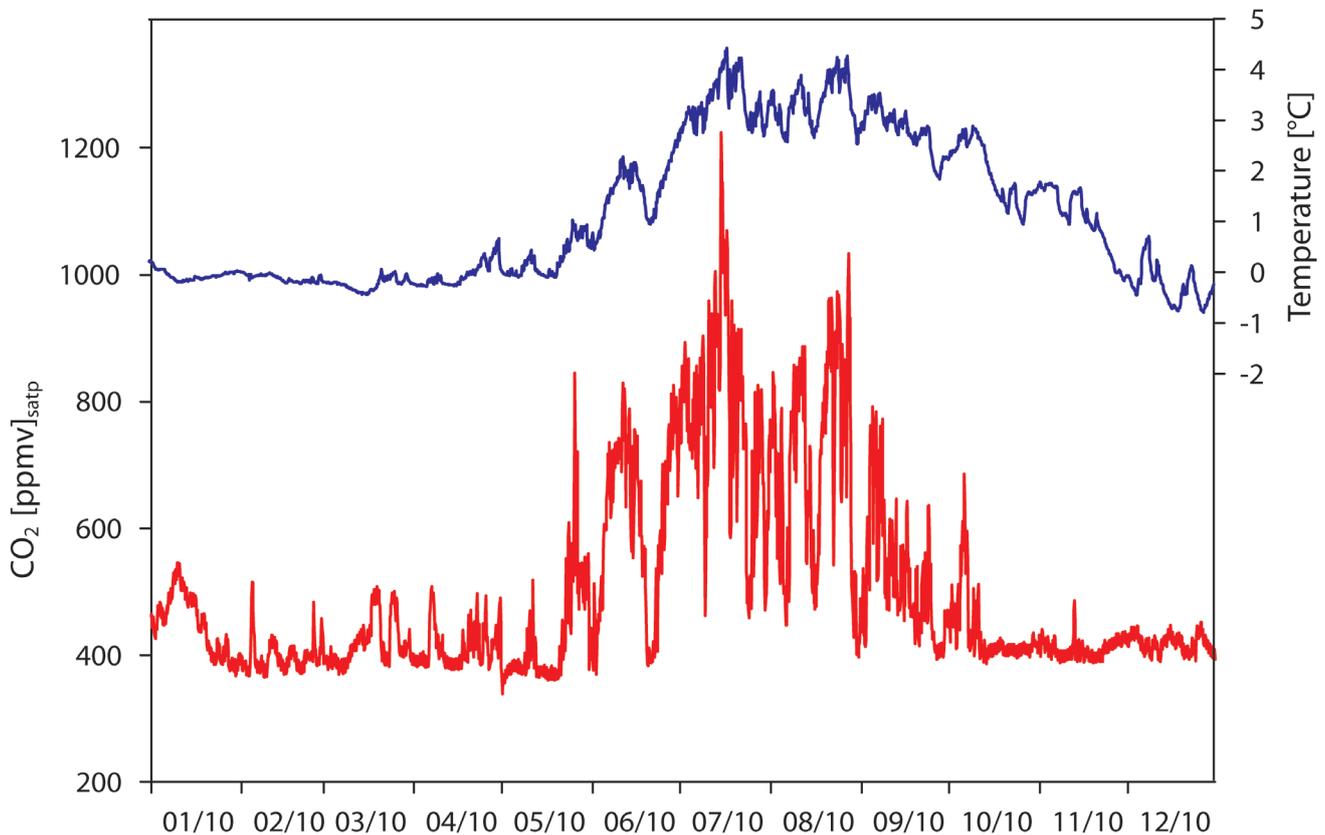


Fig. 6. Example of a [CO₂] time series of a ventilated alpine cave system obtained using a CORA instrument over a one year monitoring period.

voltage input has been improved in the latest CORA generation.

Although no long-term instrumental drift was observed during the laboratory tests, one instrument (CORA #11) showed a decreasing trend during field measurements. This linear drift is attributed to a premature aging of the sensor and was not noticed on the other instruments so far. Nonetheless, the drift can be corrected for using a calibration before and after the instrument is deployed in the field.

Because of the remarkable thermal stability of remote cave environments (e.g. Luetscher & Jeannin, 2004), significant errors associated with temperature effects are not to be expected. Changes in the ambient temperature may, however, impinge on the calibration procedure which relies on a specified reference temperature. Accordingly, attention should be paid to a stable room temperature during extended calibration work.

The carbon dioxide dilution effect due to H₂O-saturated cave atmosphere has been tested empirically by bubbling an 8000 ppmv CO₂ reference gas through distilled water until saturation was achieved (CORA #13). The resultant wet gas yielded a CO₂ value lower by 194±54 ppmv, i.e. 2.4% of the calibrated [CO₂]. Even if CO₂ dilution associated with natural fluctuations in the cave humidity generally stays within the instrumental precision, the effect has to be accounted for when studying transient ventilation regimes. Similarly, spectral interferences with isotopologues and/or other gas species as well

as band-broadening need to be considered when monitoring subtle CO₂ changes in poorly ventilated cave chambers with elevated water vapour rates. Although recent developments of laser-based instruments using wavelength-scanned cavity ring-down spectroscopy (WS-CRDS) and off axis integrated cavity output spectroscopy (ICOS) offer promising perspectives for wet sample measurements at high accuracy (Crosson, 2008; Lis et al., 2008), these technologies have not yet been tested extensively in cave environments and still require large power supplies.

For a number of applications, however, such a precise quantification of CO₂ concentrations is probably less relevant than the relative fluctuations. Results obtained during laboratory tests show that the specifications of CORA meet typical requirements of karst studies, where the spatial heterogeneity of cave air PCO₂ is frequently larger than the instrumental precision (e.g. Baldini et al., 2006).

INITIAL RESULTS FROM FIELD STUDIES

Fig. 6 shows a one-year record of CO₂ concentrations measured at 4 h intervals in the Milchbach cave system of Switzerland. The logger was installed at 140 m depth, more than 150 m from the nearest cave entrance. During the winter season, a cold air flow drawn from the lower entrances cools the cave to sub-freezing temperatures leading to occasional ice formation. Despite this rather harsh

environment data acquisition ran continuously over the monitoring period. The [CO₂] time series displays a clear seasonality with peak values measured between June and October synchronous with the maximum of soil productivity above this cave. Daily fluctuations of ca. 25 ppm are observed throughout the record pointing to a well-ventilated cave system subject to forced convection (Luetscher et al., 2011). This highly transient regime directly affects CO₂ degassing rates from cave seepage water and thus also impinges on local karstification rates. Quantified mass balances of CO₂ exchanges will be presented elsewhere.

CONCLUSIONS

The CORA monitoring device was specifically developed to monitor carbon dioxide in rough cave environments and has proven to be appropriate for the study of subsurface ventilation regimes. Its small dimensions and low power consumption make it suitable for even remote cave locations. CORA's key strengths include:

A low power consumption allowing for monitoring of at least one annual cycle;

A large measurement range adapted to the analysis of spatially and temporally variable CO₂ regimes;

Coupled temperature and pressure probes permitting atmospheric corrections;

A user-friendly software allowing to program measurement intervals and calibration parameters;

A light-weight and robust SD-Card peripheral for safe and convenient data download;

A compact and water-proof case providing protection during demanding cave expeditions.

A two-point calibration using certified reference gases allows for accurate measurements, i.e. ±50 ppmv, resp. ≤±3%vol. of the measured gas concentration.

Although no significant drift was observed, a regular re-calibration is recommended. CORA's accuracy is sufficient for a wide range of field studies ranging from ventilation to CO₂ mass balances in karst environments.

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